



Hopcroft, P., Valdes, P., O'Connor, F. M., Kaplan, J. O., & Beerling, D. J. (2017). Understanding the glacial methane cycle. *Nature Communications*, 8(14383), [14383].
<https://doi.org/10.1038/ncomms14383>

Publisher's PDF, also known as Version of record

License (if available):
CC BY

Link to published version (if available):
[10.1038/ncomms14383](https://doi.org/10.1038/ncomms14383)

[Link to publication record in Explore Bristol Research](#)
PDF-document

This is the final published version of the article (version of record). It first appeared online via Nature at <http://www.nature.com/articles/ncomms14383>. Please refer to any applicable terms of use of the publisher.

University of Bristol - Explore Bristol Research

General rights

This document is made available in accordance with publisher policies. Please cite only the published version using the reference above. Full terms of use are available:
<http://www.bristol.ac.uk/red/research-policy/pure/user-guides/ebr-terms/>

ARTICLE

Received 14 Oct 2015 | Accepted 21 Dec 2016 | Published 21 Feb 2017

DOI: 10.1038/ncomms14383

OPEN

Understanding the glacial methane cycle

Peter O. Hopcroft¹, Paul J. Valdes¹, Fiona M. O'Connor², Jed O. Kaplan³ & David J. Beerling⁴

Atmospheric methane (CH₄) varied with climate during the Quaternary, rising from a concentration of 375 p.p.b.v. during the last glacial maximum (LGM) 21,000 years ago, to 680 p.p.b.v. at the beginning of the industrial revolution. However, the causes of this increase remain unclear; proposed hypotheses rely on fluctuations in either the magnitude of CH₄ sources or CH₄ atmospheric lifetime, or both. Here we use an Earth System model to provide a comprehensive assessment of these competing hypotheses, including estimates of uncertainty. We show that in this model, the global LGM CH₄ source was reduced by 28–46%, and the lifetime increased by 2–8%, with a best-estimate LGM CH₄ concentration of 463–480 p.p.b.v. Simulating the observed LGM concentration requires a 46–49% reduction in sources, indicating that we cannot reconcile the observed amplitude. This highlights the need for better understanding of the effects of low CO₂ and cooler climate on wetlands and other natural CH₄ sources.

¹Bristol Research Initiative for the Dynamic Global Environment, School of Geographical Sciences and Cabot Institute, University of Bristol, University Road, Bristol BS8 1SS, UK. ²Met Office Hadley Centre, FitzRoy Road, Exeter EX1 3PB, UK. ³Institute of Earth Surface Dynamics, University of Lausanne, Geopolis Building, CH-1015 Lausanne, Switzerland. ⁴Department of Animal and Plant Sciences, Alfred Denny Building, University of Sheffield, Sheffield S10 2TN, UK. Correspondence and requests for materials should be addressed to P.O.H. (email: peter.hopcroft@bristol.ac.uk).

Methane (CH_4) is a potent greenhouse gas and it plays a fundamental role in atmospheric chemistry, regulating the oxidizing capacity of the troposphere and influencing surface concentrations of ozone (O_3), carbon monoxide (CO) and volatile organic compounds (VOCs) including biogenic VOCs (BVOCs). The atmospheric concentration of CH_4 has risen sharply from a concentration averaged over AD 800–1600 of 680 p.p.b.v. (parts per billion volume) characterizing the late Holocene pre-industrial era^{1,2} to 1,799 p.p.b.v. by the year 2010 (ref. 3). The increase in concentration contributed around 17% of the radiative forcing from well-mixed greenhouse gases since AD1750 (ref. 4). While the cause of increasing atmospheric CH_4 concentrations over the industrial era is unequivocally a result of anthropogenic activities, variability in CH_4 over recent decades, including a stabilization of concentrations, is not fully understood³, because of persistent uncertainties in both anthropogenic and natural processes⁵. Renewed growth in atmospheric CH_4 since 2007 is likely driven by biogenic sources, highlighting the importance of natural and/or agricultural sources of methane⁶.

The dominant natural source of CH_4 is wetlands, which constitute around 60–80% of natural emissions and around one third of the total flux to the atmosphere at present³. During the last glacial maximum (LGM), two processes may have led to a reduction in wetland CH_4 sources: (1) a reduction in wetland area, particularly in the high latitudes of both hemispheres and (2) the cooler and generally drier climate, and low atmospheric carbon dioxide (CO_2) concentrations, which led to reduced rates of methanogenesis. While large ice-sheets prevented wetland formation in North America and Northern Europe, the 120 m drop in sea-level exposed large regions of continental shelf suitable for wetland formation, in particular around Indonesia (Sunda Shelf) and between Australia and New Guinea (Arafura Shelf). Sea-level effects might offset the reduced emissions due to the other factors⁷, but most wetland schemes predict a significant reduction in emissions of between 23 and 67% relative to the pre-industrial (refs 8–11).

Reduced temperatures and humidity levels will also decrease the rate of removal of CH_4 from the atmosphere by its primary sink, that is, through reaction with the hydroxyl (OH) radical. Climate affects the rate of formation of OH by photolysis and the reaction rate of OH and CH_4 (refs 12,13). The colder and drier LGM atmosphere probably had a reduced OH burden as compared with warmer, more humid interglacial periods. A reduction in OH burden thus implies an increase in CH_4 lifetime, which could lead, paradoxically, to increased CH_4 concentrations in a LGM atmosphere. However, CH_4 is not the only chemical species that is destroyed by OH, and reductions in the emissions of other reactive trace gases could increase the amount of OH available for CH_4 oxidation. Potential expansion of tropical forests on to exposed continental shelves notwithstanding, the colder, drier climate and low atmospheric CO_2 of the LGM may have led to substantial reductions in the primary productivity of the terrestrial biosphere. This reduction could have lowered emissions of BVOCs, such as isoprene, particularly from tropical and subtropical forests⁸. A reduction in BVOC emissions could have increased OH, and hence decreased the lifetime of CH_4 . This second effect should act to offset some of the lifetime increase brought about by the overall reduction in OH burden. Thus the balance between the lifetime perturbations associated with BVOCs and physical climate is critical for understanding the LGM CH_4 lifetime, as it is for the future¹³.

Two previous studies found between 50 and 100% of the reduction in atmospheric CH_4 at the LGM was caused by a decrease in the CH_4 lifetime arising in part from reduced

BVOC emissions^{8,14}. More recently it has been proposed that the net result of the climate effects and reduced BVOC emissions on OH and hence CH_4 lifetime would be negligible, so that most if not all of the glacial reduction in CH_4 concentration should be attributed to reductions in sources¹⁵. Changes in emissions of nitrogen oxides (NO_x) from soils and lightning, and changes in photolysis rates do not alter this conclusion¹⁶. However, these past studies may have overestimated the reduction in emissions of BVOCs during the LGM, because they did not include the leaf-level stimulation of isoprene emissions caused by the low (185 ppmv) LGM atmospheric CO_2 , which potentially counters the effects from a cooler climate^{17,18}.

Here we simulate the late Holocene pre-industrial (pre-industrial in the following) and LGM atmospheric CH_4 cycles using the Earth System model HadGEM2-ES¹⁹ (see ‘Methods’ section). Here pre-industrial refers to the climate state immediately prior to Industrial Revolution around AD1750 (ref. 20). HadGEM2-ES incorporates an interactive CH_4 cycle based on schemes for dynamic vegetation and wetlands²¹ and interactive tropospheric chemistry²². We incorporate a near-complete inventory of the major natural CH_4 sources (wetlands, peatlands, biomass burning, oceans, termites and geological/hydrates) and atmospheric processes and sinks (for example, oxidation/loss by OH and soil uptake), including aspects which affect CH_4 lifetime (for example, BVOC emissions and lightning and soil NO_x production). To address aspects of stratospheric chemistry, we also include separate sensitivity tests examining O_3 photolysis and stratospheric O_3 .

Despite a lack of process-based understanding of geological, freshwater and animal CH_4 emissions, this approach is the most comprehensive yet applied to the glacial CH_4 cycle. We integrate source and sink processes together within an Earth System model, and include for the first time, estimates of peatland emissions and termite emissions, together with the CO_2 inhibition of plant isoprene emissions. We include different LGM estimates of wildfire emissions, and therefore address uncertainty in this component. We also compare the main drivers of the change in CH_4 lifetime.

We find that the Earth System model simulations do not reproduce the observed LGM CH_4 concentration, overestimating it by around 100 p.p.b.v. There is good agreement with past studies in terms of CH_4 lifetime change at the LGM, and together with significant uncertainties surrounding natural CH_4 emissions, this suggests that models of CH_4 sources are under-sensitive. If confirmed with other modelling approaches, this would have implications for understanding atmospheric CH_4 change on different timescales.

Results

Emissions calculations. The simulated LGM CH_4 emissions budget is shown in Fig. 1. This compares the predicted changes in emissions from wetlands, biomass burning, oceans, hydrates and termites, as well as bacterial uptake of methane in dry soils. Simulated wetland CH_4 emissions at the LGM are 70% of the pre-industrial (that is, a reduction from 138 to 97 $\text{TgCH}_4\text{yr}^{-1}$). Previous work shows large uncertainty in wetland emissions at the LGM, from no change⁷ to reductions similar to those we simulate^{8–10}. The wetland scheme in HadGEM2-ES accounts for changes in soil carbon as a substrate for methanogenesis^{21,23} and also the impacts of low LGM atmospheric CO_2 on the terrestrial carbon cycling. However, in common with the majority of other wetland models^{24,25}, it does not have a specific representation of hydrological and biogeochemical processes in organic soils or several other peatland-specific processes such as soil water-table variations. We addressed this issue by

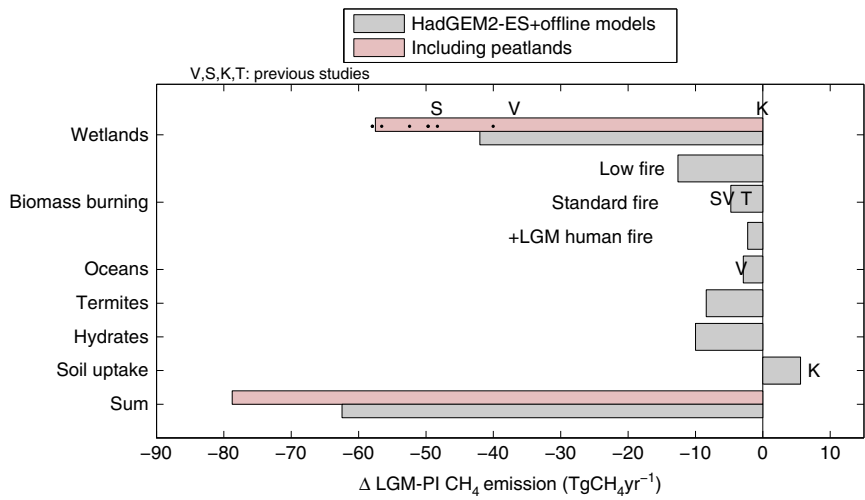


Figure 1 | Last glacial maximum minus pre-industrial CH₄ emissions. Emissions change (TgCH₄yr⁻¹) for each source term and total. Values in pink include peatland emissions in northern extra-tropics. The sum uses the standard-fire flux. Values from previous work have been scaled to the same pre-industrial fluxes as used in our study for comparison. Wetland emissions from ref. 9 (dots) based on PMIP2 coupled GCM simulations, V: Valdes *et al.*⁸, K: Kaplan *et al.*^{7,14}, T: Thonicke *et al.*²⁹ and S: Singarayer *et al.*¹⁰. More details are provided in Table 1.

Table 1 Global CH ₄ emissions diagnosed (wetlands and OH and stratospheric loss) or else prescribed within HadGEM2-ES climate-chemistry simulations.				
	PI	LGM		
	Control	Low-fire	Standard	Standard + LGM humans
Sources (TgCH ₄ yr ⁻¹)				
Wetlands	138.2	96.8	96.8	96.8
Wetlands + peatlands	138.2	80.5	80.5	80.5
Biomass burning	14.0	1.4	9.2	11.6
Oceans	15.0	12.1	12.1	12.1
Termites	20.0	11.6	11.6	11.6
Hydrates	10.0	0	0	0
SUM	197.2	121.9	129.7	132.1
Sinks (TgCH ₄ yr ⁻¹)				
Tropospheric OH loss	160	102	—	—
Soil uptake	11	6	—	—
Added stratospheric loss	21	12	—	—

Atmospheric loss terms are diagnosed within the low-fire emissions-driven simulation for consistency with the emissions listed. The LGM sum excludes the treatment of peatlands and is reduced by 16.3TgCH₄yr⁻¹ when they are considered.

performing offline simulations using the LPJ-WHyMe model of peatlands and permafrost²⁶ forced with HadGEM2-ES climate variables and the atmospheric CO₂ for each time period. With this model, the high-latitude distribution of wetland emissions is increased, possibly improving on the under-estimation of wetland regions in this area under modern climate conditions²⁷, and LGM emissions reduce from 70 to 58% of pre-industrial. This enhanced sensitivity is due to the higher proportion of the global emissions deriving from northern latitudes in the peatland model (44 TgCH₄yr⁻¹ versus 16 CH₄TgCH₄yr⁻¹ in HadGEM2-ES), where additional ice-sheets and enhanced cooling strongly reduce emissions at the LGM. Peatland processes could therefore be important for understanding wetland emissions under LGM conditions. Given the differences in sensitivity of the peatland model, it also highlights the potential for different responses in a future warmer climate.

We simulated biomass burning CH₄ emissions with LPJ-LMfire²⁸ forced offline with HadGEM2-ES climate variables

and atmospheric CO₂. This model shows a 34% reduction (from 14 to 9.2 TgCH₄yr⁻¹), consistent with the range in previous work^{8,29}. This is termed standard-fire in the following. Inclusion of an estimated human contribution to the fire at the LGM gives higher LGM fire emissions, so that the reduction in emissions at the LGM is somewhat smaller at 16% (that is, LGM emissions of 12.1 TgCH₄yr⁻¹), as shown in Fig. 1. This is termed standard + LGM humans. We also include a low-fire scenario as used in previous work¹⁶ in which all fire emissions are reduced by 90% from pre-industrial values at the LGM. This is included to bracket potential uncertainty in this term, because charcoal-based inferences imply substantial reductions in global biomass burning during the late glacial period³⁰. In this scenario, the total fire CH₄ source term reduces by a factor of 10 (from 14 to 1.4 TgCH₄yr⁻¹). The additional effects due to changes in other trace gas emissions from biomass burning on CH₄ lifetime are evaluated in the following section.

For the CH₄ flux from oceans the Guenther *et al.* BVOC scheme³¹ simulates a 20% reduction in BVOC emissions between

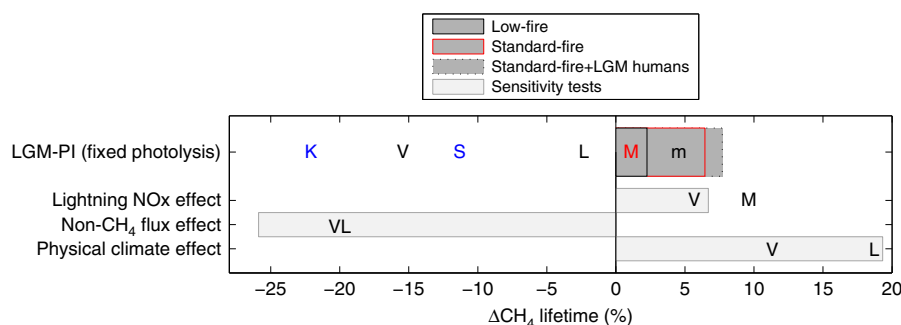


Figure 2 | Last glacial maximum atmospheric CH₄ lifetime change relative to the pre-industrial (%) and estimates of the separate contributions to this change. LGM minus pre-industrial relative CH₄ lifetime change (dark upper bar), and sensitivity experiments (perturbation minus LGM), light bars below. V: Valdes *et al.*⁸, K: Kaplan *et al.*^{7,14}, L: Levine *et al.*¹⁵, S: Singarayer *et al.*¹⁰ and Murray *et al.*¹⁶, for which only warm-LGM simulations are included and where m denotes the low-fire scenario. The effects of physical climate changes on reaction rates were not accounted for by Kaplan *et al.*¹⁴ or by Singarayer *et al.*¹⁰ shown in blue. Note for the BVOC sensitivity tests, Valdes *et al.*⁸ only changed isoprene and terpene emissions, Levine *et al.*¹⁵ also changed CO, ethane, propane and acetone; while here soil NO_x was also included. Details of all of the lifetime results are given in Table 2.

the pre-industrial and LGM. This is mostly driven by cooler sea-surface temperatures, with a small contribution from reduced surface dissolved organic carbon concentrations. To estimate emissions of CH₄ from termites, we used the observed relationships between vegetation biome type, termite species and termite CH₄ emissions per unit biomass to predict termite emissions as a function of vegetation coverage (see ‘Methods’ section). We calculated a 40% reduction at the LGM, mostly resulting from reduced tropical forest coverage. In the absence of appropriate process-based schemes for predicting hydrate and other geological CH₄ emissions, these sources are either assumed to remain unchanged at pre-industrial levels during the LGM, or alternatively were set to zero for the LGM.

Overall, the strength of the total CH₄ source is reduced by 32–46, 28–42 and 26–41% in the low-fire, standard and standard + LGM human fire scenarios, respectively, where the range depends on whether the peatland model is used and whether emissions from hydrates are reduced (Table 1).

Using the offline models for BVOC, CO and NO_x emissions (as described in ‘Methods’ section and in the Supplementary Table 1), we find relative changes that are similar to the simulated changes in CH₄ emissions. The fluxes of CO, acetone and NO_x from biomass burning are reduced by 32%, 34% and 30%, respectively, or by 21%, 15% and 7% when LGM human emissions are included. Soil NO_x emissions increase by 37%. This is partly as a result of increased LGM land area in the tropics. Transitions from dry to wetter months stimulate emissions in the soil NO_x scheme, and a general increase in the precipitation variability across the tropics stimulates NO_x emissions in the LGM simulation. The ocean CO flux is reduced by 20%. While acetone emitted from vegetation is reduced by 34%, isoprene emissions are reduced by only 19%, significantly smaller than reductions of up to 31–61%^{8,14} in previous work. This is because the stimulation of emission by low CO₂ during the LGM was not accounted for in these past studies. Lightning NO_x is modelled interactively within HadGEM2-ES (see ‘Methods’ section). There is a 25% reduction in this source, from 6.1 TgNyr⁻¹ in the pre-industrial simulation to 4.6 TgNyr⁻¹ for the LGM.

Factors influencing the glacial CH₄ lifetime. We ran HadGEM2-ES using the CH₄ source terms summarized in Fig. 1, other trace gas emissions and a soil uptake term for CH₄, all described in Table 1 and Supplementary Table 1. In this

model, the source trace gases are mostly monthly globally gridded inputs, whereas wetland CH₄ and lightning NO_x are calculated interactively, as are the main sinks of CH₄ by tropospheric OH oxidation and stratospheric loss (see Supplementary Note 1 for details).

The emissions-driven model predicts a global mean concentration of 659 ppbv for the pre-industrial close to the observed AD 800–1,600 average value of 680 p.p.b.v.¹. The lifetime of CH₄ with respect to OH is calculated using mixing ratios for CH₄ and OH up to the diagnosed model tropopause²² and is equal to 10.4 years, within the range inferred for present day³², and close to the multi-model mean lifetime 10.1 ± 1.7 years for AD1850 conditions found in the Atmospheric Chemistry and Climate Model Intercomparison Project³³. For the LGM, the lowest LGM concentration is expected in the low-fire LGM scenario, and so this is the only emissions-driven LGM scenario integrated to equilibrium. This results in a concentration of 447 p.p.b.v., which is almost 75 p.p.b.v. greater than the observed LGM concentration of 375 p.p.b.v.

Given that the emissions-driven model over-predicts the LGM CH₄ concentration, the model-derived methane lifetime will be too long given the self-feedback effect³⁴. The self-feedback effect is the positive feedback whereby any change in CH₄ concentration will influence the amount of OH available for CH₄ oxidation, which in turn further impacts on the concentration of CH₄. By running concentration-driven rather than emissions-driven simulations for the two time periods (pre-industrial and LGM), the change in CH₄ lifetime can be modelled more realistically.

In the concentration-driven simulations, the LGM lifetime increases by 2.3% relative to the pre-industrial value of 10.4 years for the low-fire scenario, and 6.5% for the standard-fire scenario as shown in Fig. 2. When the standard + LGM human fire scenario is used, the lifetime difference increases to 7.7% (see Table 2). The lifetime differences (LGM-PI) between the three fire scenarios are comparable to that reported in previous work¹⁶.

In each of the three fire cases, three main factors influence the CH₄ lifetime: reduced atmospheric temperatures and humidity levels, changes in surface emission rates of CO, NO_x, isoprene and acetone, and a reduction in the lightning generation of NO_x as simulated within HadGEM2-ES. Physical climate changes can be summarized by a global mean cooling of 4.8 °C and a global reduction in water vapour of 24% at the surface, with similar humidity reductions throughout the atmospheric column. The other tropospheric O₃ precursor emission

Table 2 | Global summary of the simulated CH₄ lifetime with respect to OH ($\tau_{\text{CH}_4,\text{OH}}$).

	Pre-Ind	LGM		
	Control	Low-fire	Standard	+ LGM humans
Emissions-driven				
Offline photolysis	10.4	11.1 (+ 6.5%)	—	—
Offline photolysis, PI lightning NO _x	—	9.9	—	—
Offline photolysis, PI non-CH ₄ surf. fluxes	—	13.5	—	—
Offline photolysis, PI climate	—	8.6	—	—
Concentration-driven				
Offline photolysis	10.5	10.7 (+ 2.3%)	11.2 (+ 6.5%)	11.3 (7.7%)
Fast-J photolysis	7.3	7.3 (− 0.2%)	—	—
Fast-J photolysis, + 3% strat. O ₃	—	7.4 (+ 1.7%)	—	—
Fast-J photolysis, Murray <i>et al.</i> ¹⁶ strat. O ₃	—	7.4 (+ 2.1%)	—	—

For the three fire-emissions scenarios: (i) low-fire, (ii) standard (based on LPI-LMfire) and (iii) standard with LGM human-induced fires at the LGM. The percentage change relative to the respective pre-industrial control simulation is given in brackets.

changes are summarized above, with full details in the ‘Methods’ section.

We quantified the relative influences on the CH₄ lifetime with three additional LGM sensitivity tests with the low-fire interactive CH₄ (that is, emissions-driven) set-up. The three additional emissions-driven simulations were LGM but with pre-industrial tropospheric O₃ precursor emissions excluding methane, LGM but with pre-industrial lightning NO_x generation rates and LGM but with pre-industrial physical climate. We find that LGM non-CH₄ emissions cause a 26% decrease in lifetime at the LGM. The LGM lightning reduction is responsible for a lifetime increase of 7%, while the physical climate changes increase the lifetime by 19%. These components of the total lifetime change are shown in Fig. 2. They are comparable with previous work^{8,15,16} though the stronger cooling in HadGEM2-ES compared with the previous model simulations, is likely responsible for the larger climate term.

For computational efficiency HadGEM2-ES uses pre-computed photolysis rates from an offline 2D model³⁵. We also evaluate a model version including an interactive photolysis scheme³⁶, which responds dynamically to changes in clouds and aerosols within HadGEM2-ES. This allows the effects from changes in stratospheric O₃ concentrations on incoming radiation incident on the troposphere, and hence on the production of OH, to be included. Switching from prescribed photolysis rates to interactive photolysis rates together with a prescribed increase in stratospheric ozone¹⁶ has a very small impact on the methane lifetime, as detailed in the Supplementary Note 2.

CH₄ concentration sensitivity to source and lifetime changes.

We now combine the model predicted estimates of source and lifetime changes to derive a possible range of concentration predictions for the LGM. Because of the high computational cost of the coupled HadGEM2-ES model, we use a simplified mass balance formulation (described in ‘Methods’ section) to calculate the concentration as a function of the global CH₄ source and the CH₄ lifetime, and assuming steady state conditions in the two time periods, though this may not hold for all of the late Holocene pre-industrial period³⁷. In terms of the sources, we incorporate all of the different source possibilities, including peatlands, changes to hydrates and the three fire scenarios described previously. For lifetime, we apply three different possible changes which correspond to the simulated changes in CH₄ lifetime in response to the three fire scenarios. We consider changes in sources and sinks separately and then combinations of the two, to give a summary of how the

LGM CH₄ budget depends on the processes analysed. This analysis follows a similar approach recently presented for glacial CO₂ (ref. 38).

The individual contributions of the source and sink terms considered in this work are shown for comparison in Fig. 3. Our analysis shows that the largest changes in CH₄ concentrations in absolute terms arise from changes in either the wetland source or tropospheric O₃ precursors including BVOCs, justifying their focus in previous work. However, the combined contribution from other terms is significant, especially for the sinks. In past work, only relatively few of these terms have been quantified. The most studied aspect is the LGM wetland source for which previous work shows a significant spread, which encompasses the difference between the wetland and peatland models used here.

With a constant CH₄ lifetime, the wetland and low-fire sources are the largest terms in the PI-LGM amplitude. The other fire terms are significantly smaller. Changes in the lifetime given a constant source are generally larger contributors than many of the individual CH₄ source terms. The net effect of the lifetime change is small though, because it comprises a combination of several large values of opposing signs.

A successful simulation of the LGM to pre-industrial concentration change must reproduce both the pre-industrial concentration value and the change in concentration between the pre-industrial and LGM, which is 305 p.p.b.v., as shown in Fig. 3. This ice-core inferred amplitude is associated with negligible uncertainty, so we concentrate on model uncertainty in the following. The first criterion is nearly satisfied in this model, as the pre-industrial concentration is 660 p.p.b.v. about 20 p.p.b.v. lower than observed for the period 800–1,600AD. However, a positive trend during this period may result from pre-industrial era anthropogenic activities¹. Our simulations do not include any anthropogenic sources, and so the 20 p.p.b.v. under-estimate for the pre-industrial in the model is consistent with such a human contribution during this time period. However, uncertainties in each of the source and sink terms as further discussed below, preclude confidence in the attribution of the 20 p.p.b.v. difference between the simulation and observations.

We now assess whether imposing the predicted changes in CH₄ sources and CH₄ lifetime (τ) can reproduce the second of the ice-core derived criteria described above. The net effect is estimated by combining for each of the three fire scenarios, the four source estimates (hydrates changing or not, and including peatlands or not) with the respective fire-scenario lifetime change. This gives four concentrations for each of the three scenarios.

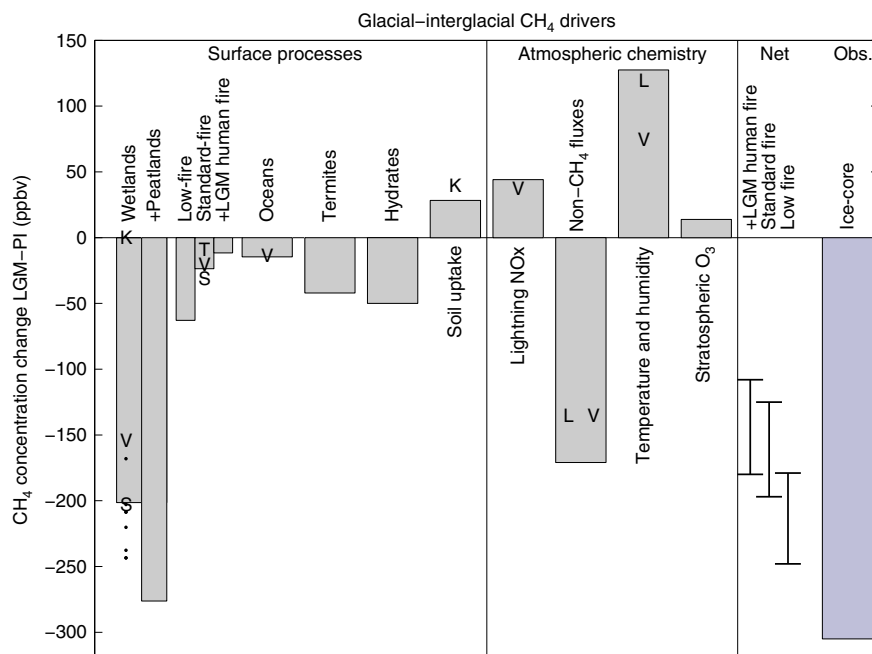


Figure 3 | Summary of the contributions to the glacial to inter-glacial CH_4 budget as modelled in this work and in previous studies. V: Valdes *et al.*⁸, K: Kaplan *et al.*^{7,14}, L: Levine *et al.*¹⁵, T: Thonicke *et al.*²⁹ and S: Singarayer *et al.*¹⁰. The contributions from individual sources are calculated assuming no change in lifetime, while the contributions from sinks are calculated with constant source terms. The total concentration changes take account of both source and sink changes. The separate contributions to the change in chemical lifetime of CH_4 are quantified for the low-fire scenario here. The non- CH_4 term will be smaller for the standard fire and standard + LGM humans fire scenarios. The uncertainty range in each of the 'Net' values, encompasses whether or not peatlands are considered. Details of these calculations are given in Supplementary Table 2.

In the three fire scenarios, the total CH_4 LGM source reduction required to give the observed LGM concentration can be calculated with the mass balance formulation (see 'Methods' section). This results in required source reductions of 46%, 48% or 49% for the low-fire, standard-fire and standard + LGM human fire scenarios, respectively. This range is at the extreme end of our source predictions, which showed maximum reductions of 46%, 42% and 41% for the same three fire scenarios, respectively. These reductions are only 41%, 36% and 34% if the hydrate term is not reduced at the LGM.

The estimated concentration change assuming no change in hydrate emissions is shown in Fig. 3 (and details of the calculations are listed in Supplementary Table 2). For either of the process-based fire estimates (standard or with LGM humans), the resultant CH_4 concentration predictions all fall short of the observed amplitude at 125–196 p.p.b.v. and 108–181 p.p.b.v. for the standard fire and standard fire with LGM human scenarios, respectively. The range depends on whether peatland model is used or not, with larger changes when the peatland model is used. The low-fire model has both the largest decrease in the CH_4 source, and the smallest increase in lifetime at the LGM. The modelled PI-LGM concentration change is nearly consistent with the observations in the low-fire scenario when taking the very upper range of the predicted 179–248 p.p.b.v. change.

Conditional on the pre-industrial hydrates flux of $10 \text{ TgCH}_4\text{yr}^{-1}$ reducing to zero at the LGM then the predicted concentration changes are larger, at 169–240, 152–224 and 221–290 p.p.b.v. for the standard fire, standard with LGM humans and low-fire scenarios, respectively. We note that the low-fire scenario is unlikely to be consistent with the isotopic record of $\delta^{13}\text{CH}_4$, because biomass burning is a source of ^{13}C enriched CH_4 , and a significant reduction in biomass burning

emissions during the LGM will lead to a depletion in atmospheric $\delta^{13}\text{CH}_4$ relative to the pre-industrial. This is opposite to the observed 6‰ (permille) enrichment³⁹. For a best estimate, we therefore leave out the low-fire scenario, assume hydrate emissions did not change at the LGM (there is little evidence to support a complete shutdown of this source) and include the peatland model. This leads to a calculated LGM concentration of between 463–480 p.p.b.v. for the standard and standard with LGM human fire scenarios. This is 89–105 p.p.b.v. higher than observed for the LGM.

Uncertainty in the source mix for the pre-industrial describes the incomplete knowledge of the relative contributions of different natural (and anthropogenic) CH_4 sources in the present day and pre-industrial³. This is especially important for non-wetland sources such as freshwater, wild animals and geological sources³, (see Supplementary Table 3 and Supplementary Note 3). The range of concentrations that arises from this source of uncertainty is assessed in Supplementary Note 3. In these scenarios, we do not include any change in the geological or wild animal sources, but we assume that freshwater emissions scale with wetland emissions. Future work will need to consider recent process-based models of freshwater emissions (ref. 40) to address this assumption. We calculate a maximum difference of 30 p.p.b.v. between three different predictions of the LGM CH_4 concentration. This is therefore comparable to the range induced by the LGM lifetime changes. However, it is not currently possible to robustly discern between the different source scenarios.

Discussion

Several aspects of the current understanding of the global CH_4 cycle are subject to uncertainty. We highlight some key

areas that are relevant to the LGM CH₄ budget and relate them to the results summarized in Fig. 3.

We demonstrate that inclusion of peatland processes is essential for realistically capturing the climatic response. Several other processes are potentially important but require further detailed investigation in this context. These include wetland-specific plant functional types^{26,41}, seasonal flooding by river systems⁴¹, ground water controls on wetland formation⁴², nutrient status⁴³ and CO₂ effects on CH₄ emissions from different types of wetlands⁴⁴. Future work is required to evaluate the relative importance of these processes, though new observations of the CH₄ cycle may well be required²⁵.

Fire is an important natural source of CH₄, BVOCs, CO and NO_x. Based on charcoal assemblages, wildfires were thought to be very strongly reduced during the glacial period³⁰, although there are relatively few sites covering the LGM. Model predictions are less extreme, showing only a 20–40% reduction in emissions. Because fire affects both the CH₄ source and the lifetime, we explored the potential influence of biomass burning CH₄, by including a low-fire scenario in our simulations, but further work is required. The estimated influence of humans on fires during the LGM is non-negligible, but the associated uncertainties are difficult to quantify at present.

Modern day CH₄ emissions derived from wild animals, geological sources and freshwater bodies have high relative uncertainty of ~100% around central estimates³. Our additional source scenarios (Supplementary Table 3) account for some of this uncertainty by including two alternative estimates for the relative magnitude of the pre-industrial CH₄ sources. We find that the composition of natural CH₄ sources that we specify is an important factor for the LGM budget, because large sources that are not directly affected by climate (for example, wild animals or geological) would pose problems for satisfying the LGM CH₄ budget. Our results imply that large geological and wild animal sources of CH₄ are not compatible with the presented LGM CH₄ budget.

The reduced lightning NO_x flux plays a major role in the chemical lifetime of CH₄ at the LGM, increasing it by around 10% and hence offsetting much of the decrease in sources. However, the response of atmospheric convection and lightning to global climate change is highly uncertain⁴⁵. HadGEM2-ES reproduces some aspects of the enhanced dust cycle for the LGM⁴⁶, as inferred from sedimentary archives⁴⁷. However, a hypothesized link between aerosol particles, cloud microphysics and lightning frequency⁴⁸ is not parameterized, but could conceivably have modulated lightning NO_x emissions during the LGM.

Large uncertainties surround the prediction of BVOC emissions as a function of climate and ambient CO₂ (ref. 49). One recent study predicted a 15% increase in isoprene emissions at the LGM¹⁸ with an empirically defined parameterization of the CO₂ suppression effect. This contrasts with the 19% reduction found here. A 15% increase in BVOC emissions would mean even larger CH₄ emissions reductions are required during the LGM relative to the pre-industrial. These results imply there are major uncertainties in the sensitivity of isoprene emissions to climate state. Recent progress in process-based understanding may provide new insight into BVOC emissions responses to past climates^{49,50} as it relates synthesis of isoprene to excess electron transport over that required for photosynthesis within the leaf. The advantage of this is that it removes the need for separate empirical parameterisations of the influence of different environmental controls of BVOC emissions as have been used to date.

The short-wave cloud forcing for the LGM-PI has been shown for this model previously⁵¹. Despite marked changes over

tropical land areas, we find a relatively limited effect via photolysis rates on the LGM CH₄ lifetime in HadGEM2-ES. The new version of the Hadley Centre model, HadGEM3 includes a prognostic as opposed to diagnostic cloud scheme⁵² and this shows much higher cloud cover fractions than HadGEM2-ES. This will likely alter the predicted response of photolysis rates when this newer physical model is coupled with an interactive chemistry scheme. There is also uncertainty in the radiative forcing from stratospheric ozone in response to elevated CO₂ in climate chemistry model simulations^{53,54}, that may be important for understanding model dependency in oxidizing capacity related to stratospheric chemistry processes.

In conclusion, Quaternary fluctuations in greenhouse gas concentrations can serve as fundamental tests for understanding of the Earth System. The CH₄ cycle is particularly interesting because it is also highly sensitive to abrupt climate change with rapid rises of up to 200 p.p.b.v. over a few decades during Dansgaard–Oeschger events of the glacial period⁵⁵. CH₄ is also an important greenhouse gas, and as such it is imperative to improve our understanding of the feedbacks between climate and natural CH₄ emissions, and the LGM is a potentially valuable example, because it is likely that emissions were substantially reduced.

We used an Earth System model and a series of offline models of trace gas emissions from natural sources to quantify the primary controls on the atmospheric CH₄ cycle during the last glacial maximum relative to the pre-industrial. Our comprehensive modelling framework incorporates several relatively untested processes, including peatland CH₄ emissions, CO₂ inhibition of isoprene and a process-based representation of natural and anthropogenic biomass burning. However, the modelled best-estimate LGM CH₄ concentration is between 463 and 480 p.p.b.v., which significantly overestimates the observed concentration of 375 p.p.b.v.

There is relatively good agreement between previous work and results reported here for the relative influence of climate, BVOCs and lightning on the LGM CH₄ lifetime. There is also relatively limited lifetime change simulated for the LGM, and this implicates a major reduction of CH₄ sources during the LGM^{15,16}. Uncertainty in CH₄ sources also appears to outweigh that in the sink processes, because of the large inter-model uncertainty in wetland CH₄ schemes²⁵ and the large uncertainty surrounding estimates for other natural sources³. This is somewhat different to the situation for the historical and future studies where uncertainty in the lifetime is important.

Current CH₄ emission models appear to underestimate the change needed to balance the LGM budget. This could be attributed either to incorrect parameterization of processes in existing models, or to missing processes, as outlined above. An enhanced sensitivity of CH₄ sources would be important, as it could alter our understanding of climatically driven variations in atmospheric CH₄ concentration, such as during rapid rises of the last glacial⁵⁵ and possibly during recent decades and the future.

Further work is required to systematically identify whether factors including groundwater control on wetland formation, wetland-specific carbon cycling or others listed, will change the environmental sensitivity of these aspects of the CH₄ cycle. Future work could benefit considerably from a formalized probabilistic analysis of this problem, following a similar approach taken for aerosols⁵⁶. This would allow uncertainties in all of the model components to be propagated through to final predictions, allowing an objective assessment of whether the observations are reconciled by the model, and potentially highlighting key uncertainties in a more systematic and objective manner.

Methods

Earth system model. We use a coupled atmosphere-vegetation-chemistry model HadGEM2-ES¹⁹ to analyse the LGM to pre-industrial atmospheric CH₄ increase. HadGEM2-ES is a comprehensive Earth System model, and consists of a 3D ocean general circulation model (GCM) coupled to a 3D atmospheric GCM, although here we utilize an atmosphere-only configuration. The atmosphere has a horizontal resolution of $1.875^\circ \times 1.25^\circ$ with 38 unequally spaced vertical levels reaching to over 39 km. The model simulates interactive dynamic vegetation⁵⁷, dynamic wetlands and wetland CH₄ emissions²¹. It includes a tropospheric chemistry scheme from the United Kingdom Chemistry and Aerosols (UKCA) model²², which can be coupled with the dynamic wetland emissions scheme. The comprehensive aerosol scheme includes mineral dust (using six tracers) which is coupled to the dynamic vegetation and land surface schemes¹⁹. HadGEM2-ES does not include stratospheric halogen chemistry²², and so ozone mixing ratios from three levels above the dynamically defined tropopause are prescribed based on an observationally based zonally averaged climatology⁵⁸.

Here HadGEM2-ES is coupled to the TropIso configuration of UKCA²², which includes isoprene chemistry missing in the standard CMIP5 version of HadGEM2-ES^{19,58}. TropIso employs the lumped Mainz Isoprene Mechanism^{22,59}, but is otherwise identical to that implemented in HadGEM2-ES for CMIP5 (ref. 58). The model includes 54 chemical species of which 37 are model tracers. It simulates 35 photolytic reactions, 114 bimolecular reactions and 15 uni- and termolecular reactions²². Photolysis rates are pre-calculated using an offline 2D model³⁵ with a two-stream approach. This does not take account of changes in clouds, stratospheric O₃ and aerosols, but gives diurnal and seasonal variability. Lightning nitrogen oxide (NO_x) emissions are calculated within HadGEM2-ES²² using a well-established scheme^{60,61}. This has been re-tuned in comparison with the version quoted in ref. 22, and gives a pre-industrial global source strength of 6.1 TgNyr^{-1} .

The pre-industrial and LGM simulations follow the Paleoclimate Model Intercomparison Project 2 (PMIP2) protocols as used in previous work⁵¹. Land use is not included in either time-period and potential vegetation is simulated dynamically. For the LGM, the model shows a mean tropical cooling of 2.7°C (consistent with the range of $1.6\text{--}3.2^\circ\text{C}$ in a recent model intercomparison⁶²), and an expansion of sea-ice coverage from 5.5 to 7.5% averaged globally. The LGM simulation without interactive chemistry has been evaluated against LGM reconstructions of terrestrial biome distributions, surface temperatures⁵¹ and dust deposition rates⁴⁶ and reproduces enhanced cooling over land and toward the poles, a reduction in extra-tropical forest coverage and increased dust deposition rates seen in reconstructions. When the chemistry model is employed atmospheric CH₄ is either allowed to evolve freely (emissions-driven), or is prescribed at the surface (concentration-driven) with values of 675 and 375 p.p.b.v. for the pre-industrial and LGM, respectively.

Trace gas emissions and soil CH₄ uptake calculations. The chemistry scheme is driven with monthly varying trace gas emissions from wetlands (CH₄), biomass burning (CH₄, CO and NO_x), vegetation (isoprene and acetone), soil NO_x, oceans (CH₄ and CO) and termite and hydrate CH₄, as well as soil uptake of CH₄. A summary of the sources and sinks considered is given in Supplementary Table 4.

Wetland CH₄ emissions are computed within HadGEM2-ES at each model timestep¹⁹ with a TOPMODEL scheme^{21,63,64} here configured with a recently derived high-resolution global topographic index data set⁶⁵, which describes the propensity for subgrid areas of soil moisture saturation. Wetland emissions are dependent on soil carbon, soil temperature and the simulated wetland fraction. Soil carbon in HadGEM2-ES is represented by four pools with different turnover rates. The four pools are incremented by litterfall from the dynamic vegetation scheme and are depleted by soil aerobic respiration which is a function of soil temperature and moisture²³.

Due to the 120 m lowering of sea-level at the LGM, exposed areas of continental shelf can support new wetland areas. The unknown topographic index for these points (currently sub-sea) were calculated using the sub-grid scale topographic variability derived from high-resolution ETOPO1 data set⁶⁶, which has global (land and ocean) coverage. A linear best fit between the observed topographic index field⁶⁵ and the log of orographic roughness performed at 10 arc-minute resolution was used to calculate the topographic index values for exposed areas of continental shelf⁵¹.

Biomass burning trace gas emissions are applied in HadGEM2-ES following the distribution and magnitude from ref. 67 for the year 1850. The change at the LGM is calculated from simulations with the LPJ-LMfire model²⁸. LPJ-LMfire accounts for natural ignition, multi-day burning and coalescence of wildfires. It includes a precipitation dependency of burning versus smouldering, and a representation of human influences. The model was driven with monthly climatologies from HadGEM2-ES combined in an anomaly mode with detrended observational climatologies listed in Table 3 of ref. 28. Land use for the pre-industrial is taken for the year 1770. Two LGM scenarios were run, one without a human influence and a second including a parameterization of the fire activities of hunter-gatherers estimated for this time period⁶⁸.

Termite CH₄ emissions were calculated on a monthly basis from simulations with the BIOME4 model driven with HadGEM2-ES pre-industrial and LGM climatologies (in an anomaly mode at 0.5° resolution, similar to that

described for LPJ-LMfire above). The resultant vegetation distribution was then used with the observed termite biomass per unit area of forest versus grasslands and the observed CH₄ emissions per unit biomass⁶⁹ to estimate the global emissions. This results in a global pre-industrial source close to $20 \text{ TgCH}_4\text{yr}^{-1}$ from independent estimates³.

Soil uptake was simulated using the model of ref. 70. This used monthly climatologies of soil temperature (at 10 cm), frozen soil fraction and wetland fraction (where no uptake is allowed) from HadGEM2-ES. Soil types were derived from the Harmonized World Soil Database v1.2 (ref. 71) and were extrapolated to the LGM land-sea-mask.

Acetone and isoprene emissions from plants were simulated using the Joint UK Land Environment Simulator (JULES) version 2.2 (refs 23,64), which incorporates a biogenic emissions model^{72,73}. JULES was forced with 3-hourly output from HadGEM2-ES simulations of the pre-industrial and LGM.

The soil NO_x model is a semi-empirical scheme⁷⁴, which has been used previously in this context⁸. The ocean BVOC model³¹ was used in the same manner as previous studies⁸, except here we also account for changes in surface ocean dissolved organic carbon between the pre-industrial and LGM by using output from IPSL-CM5A-LR Earth System Model simulations of the pre-industrial and LGM⁷⁵ as archived in the CMIP5 database.

Peatlands CH₄ emissions are simulated using LPJ-WHyMe²⁶, a process-based model of peatland and permafrost processes that includes two peatland plant functional types, vertically resolved production, oxidation and transport of CH₄ in soil and water table dynamics. The peatland area is prescribed from HadGEM2-ES as the fractional coverage with saturated soil conditions following ref. 76. This results in a pre-industrial peatland CH₄ flux of $44 \text{ TgCH}_4\text{yr}^{-1}$, comparable to other estimates⁷⁷. This is nearly three times larger than the $16 \text{ TgCH}_4\text{yr}^{-1}$ simulated with the HadGEM2-ES wetland scheme for an identical wetland area. For consistency when including the peatland fluxes in budget calculations, the non-peatland wetland CH₄ flux is scaled so that the global pre-industrial wetland plus peatland CH₄ flux is unchanged.

The remaining trace gas emissions do not change between pre-industrial and LGM simulations, and follow the distributions used in HadGEM2-ES for the CMIP5 pre-industrial simulation⁵⁸, representative of the year 1860. These emissions are C₂H₆ (ethane), C₃H₈ (propane), HCHO (formaldehyde) and MeCHO (acetaldehyde), where Me = CH₃. The overall non-CH₄ trace gas emission budget is summarized in Supplementary Table 1.

Model simulation set-up. The model simulations consist of three phases (see Supplementary Fig. 1). In phase 1, pre-industrial and LGM HadGEM2-ES simulations were carried out with the atmospheric chemistry deactivated, as reported previously⁵¹. In these HadGEM2-ES simulations, the soil carbon and vegetation are in equilibrium⁵¹. The monthly climatologies were then used to force each of the offline trace gas emissions models and the soil CH₄ uptake model described above.

In phase 2, the physical atmosphere and land surface fields in the coupled chemistry-climate HadGEM2-ES simulations were initialized from the final state of these phase 1 HadGEM2-ES simulations. The atmospheric chemistry trace gases were initialized with fields from a separate pre-industrial simulation with HadGEM2-ES with tropospheric chemistry activated. The climate-chemistry model was then spun up for 100 years using the trace gas emissions and soil CH₄ uptake calculated with offline models in phase 1. The last 30 years of each chemistry-climate simulation were averaged for analysis.

In phase 3, further sensitivity simulations were branched off from year 50 of the phase 2 simulations. In the sensitivity simulations either (i) lightning, (ii) CO, NO_x and BVOCs, or (iii) climate follow pre-industrial conditions but the rest of the model sees LGM conditions. Simulations (i) and (ii) were branched off from the emissions-driven LGM simulation and simulation (iii) was branched off from the pre-industrial simulation. These sensitivity simulations are at least 50 years long, and trends of less than $10 \text{ p.p.b.v. per } 100 \text{ years}$ in global mean surface CH₄ mixing ratio are taken as steady state, which is satisfied in each case (see Supplementary Fig. 2 and Supplementary Note 1). Prescribed surface CH₄ mixing ratio (that is, concentration-driven) simulations were also branched off from the dynamic emissions-driven runs in phase 2 and are 30 years long for the offline photolysis and 20 years long when the photolysis rates are interactively calculated within HadGEM2-ES (see Supplementary Note 1).

Offline CH₄ concentration calculation. Because of the computational expense of HadGEM2-ES, we use a simplified mass balance calculation of the atmospheric concentration to test the sensitivity to different combinations of total CH₄ emissions and lifetime changes between the pre-industrial and LGM. For this, we employed the formulation $\frac{dB}{dt} = S - B/\tau$, where B is the tropospheric CH₄ burden in Tg, S is the global CH₄ source ($\text{TgCH}_4\text{yr}^{-1}$) and τ is the tropospheric CH₄ lifetime in years, and assuming steady state in each time period, that is, $\frac{dB}{dt} = 0$. B is converted to a volume concentration using a conversion factor k which can vary as a function of tropopause height and the distribution of atmospheric CH₄. HadGEM2-ES predicts that k reduces by 2.7% at the LGM, so we include this reduction from $k = 2.6 \text{ Tg p.p.b.v.}^{-1}$ for the pre-industrial. τ is set to 10.4 years for the pre-industrial as simulated in HadGEM2-ES. The atmospheric

concentration for the PI is 660 p.p.b.v. This formulation reproduces the emissions-driven LGM concentration simulated by HadGEM2-ES to within 2%.

A self-feedback factor of $F = 1.26$ was calculated from the emissions-driven and concentration-driven lifetime and concentration values (Table 2) following equations 1 and 2 of ref. 78. This is used to augment the LGM $[\text{CH}_4]$ for the source-induced $[\text{CH}_4]$ changes, which otherwise do not include this feedback. Changes in the net source and lifetime are imposed based on the offline trace gas emission simulations and the coupled climate-chemistry simulations (Tables 1 and 2). Hence the self-feedback factor for CH_4 lifetime is implicitly included. Details of these calculations are given in Supplementary Table 2 and examples are given in Supplementary Note 4.

The required source to reproduce the observed LGM CH_4 concentration of 375 p.p.b.v. can be calculated by using the lifetime changes for the three scenarios and the equation above. In this case, the constant k is reduced by 2.7% for the LGM calculation. This results in S values for the LGM of 89, 86 and 85 $\text{TgCH}_4\text{yr}^{-1}$ for the low-fire, standard-fire and standard + LGM human scenarios, respectively. Relative to a net source of 165 $\text{TgCH}_4\text{yr}^{-1}$ for the pre-industrial (Table 1), these values represent reductions of 46%, 48% and 49%, respectively.

Data availability. The source code for HadGEM2-ES is part of the UK Met Office Unified Model and is subject to Crown Copyright. The Met Office Unified Model is available for use under licence. For more information see <http://www.metoffice.gov.uk/research/collaboration>.

The Earth System model variables analysed in the study are available from www.bridge.bris.ac.uk/resources/simulations.

References

- Mitchell, L. *et al.* Constraints on the late Holocene anthropogenic contribution to the atmospheric methane budget. *Science* **342**, 964–966 (2013).
- MacFarling Meure, C. *et al.* Law dome CO_2 , CH_4 and N_2O ice core records extended to 2000 years BP. *Geophys. Res. Lett.* **33**, L14810 (2006).
- Kirschke, S. *et al.* Three decades of global methane sources and sinks. *Nat. Geosci.* **6**, 813–822 (2013).
- Myhre, G. *et al.* in *Climate Change 2013: The Physical Science Basis: Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change* (eds Stocker, T. *et al.*) (Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2013).
- Nisbet, E., Dlugokencky, E. & Bousquet, P. Methane on the Rise—Again. *Science* **343**, 493–494 (2014).
- Schaefer, H. *et al.* A 21st century shift from fossil-fuel to biogenic methane emissions indicated by $^{13}\text{CH}_4$. *Science* **352**, 80–84 (2016).
- Kaplan, J. Wetlands at the last glacial maximum: distribution and methane emissions. *Geophys. Res. Lett.* **29**, 1079 (2002).
- Valdes, P., Beerling, D. & Johnson, C. The ice age methane budget. *Geophys. Res. Lett.* **32**, L02704 (2005).
- Weber, S., Drury, A., Toonen, W. & van Wee, M. Wetland methane emissions during the last glacial maximum estimated from PMIP2 simulations: climate, vegetation, and geographic controls. *J. Geophys. Res.* **115**, D06111 (2010).
- Singarayer, J., Valdes, P., Friedlingstein, P., Nelson, S. & Beerling, D. Late Holocene methane rise caused by orbitally controlled increase in tropical sources. *Nature* **470**, 82–85 (2011).
- Ringeval, B. *et al.* Response of methane emissions from wetlands to the Last Glacial Maximum and an idealised Dansgaard-Oeschger event: insights from two models of different complexity. *Clim. Past* **9**, 149–171 (2013).
- Johnson, C., Stevenson, D., Collins, W. & Derwent, R. Role of climate feedback on methane and ozone studied with a coupled Ocean-Atmosphere-Chemistry model. *Geophys. Res. Lett.* **28**, 1723–1726 (2001).
- O'Connor, F. *et al.* Possible role of wetlands, permafrost and methane hydrates in the methane cycle under future climate change: a review. *Rev. Geophys.* **48**, RG4005 (2010).
- Kaplan, J., Folberth, G. & Hauglustaine, D. Role of methane and biogenic volatile organic compound sources in the late glacial and Holocene fluctuations of atmospheric methane concentrations. *Glob. Biogeochem. Cyc.* **20**, GB2016 (2006).
- Levine, J. *et al.* Reconciling the changes in atmospheric methane sources and sinks between the last glacial maximum and the pre-industrial era. *Geophys. Res. Lett.* **38**, L23804 (2011).
- Murray, L. *et al.* Factors controlling variability in the oxidative capacity of the troposphere since the Last glacial maximum. *Atmos. Chem. Phys.* **14**, 3589–3622 (2014).
- Possell, M., Hewitt, C. & Beerling, D. The effects of glacial atmospheric CO_2 concentrations and climate on isoprene emissions by vascular plants. *Global Change Biol.* **1**, 60–69 (2005).
- Achakulwisut, P. *et al.* Uncertainties in isoprene photochemistry and emissions: implications for the oxidative capacity of past and present atmospheres and for climate forcing agents. *Atmos. Chem. Phys.* **15**, 7977–7998 (2015).
- Collins, W. *et al.* Development and evaluation of an Earth-system model—HadGEM2. *Geosci. Model Dev.* **4**, 1051–1075 (2011).
- Braconnot, P. *et al.* Results of PMIP2 coupled simulations of the Mid-Holocene and Last Glacial Maximum Part 1: experiments and large-scale features. *Clim. Past* **3**, 261–277 (2007).
- Gedney, N., Cox, P. & Huntingford, C. Climate feedback from wetland methane emissions. *Geophys. Res. Lett.* **31**, L20503 (2004).
- O'Connor, F. *et al.* Evaluation of the new UKCA climate-composition model—Part 2: The Troposphere. *Geosci. Model Dev.* **41**–91 (2014).
- Clark, D. *et al.* The Joint UK Land Environment Simulator (JULES), model description—Part 2: carbon fluxes and vegetation dynamics. *Geosci. Model Dev.* **4**, 701–722 (2011).
- Wania, R. *et al.* Present state of global wetland extent and wetland methane modelling: Methodology of a model intercomparison project (WETCHIMP). *Geosci. Model Dev.* **6**, 617–641 (2013).
- Melton, J. *et al.* Present state of global wetland and wetland CH_4 modelling: conclusions from a model intercomparison project (WETCHIMP). *Biogeosciences* **10**, 753–788 (2013).
- Wania, R., Ross, I. & Prentice, I. C. Implementation and evaluation of a new methane model within a dynamic global vegetation model: LPJ-WHyMe v1.3.1. *Geosci. Model Dev.* **3**, 565–584 (2010).
- Hayman, G. *et al.* Comparison of the HadGEM2 climate-chemistry model against *in-situ* and SCIAMACHY atmospheric methane data. *Atmos. Chem. Phys.* **14**, 13257–13280 (2014).
- Pfeiffer, M., Spessa, A. & Kaplan, J. A model for global biomass burning in preindustrial time: LPJ-LMfire (v1.0). *Geosci. Model Dev.* **6**, 643–685 (2013).
- Thonicke, K., Prentice, I. & Hewitt, C. Modeling glacial-interglacial changes in global fire regimes and trace gas emissions. *Glob. Biogeochem. Cyc.* **19**, GB3008 (2005).
- Power, M. *et al.* Changes in fire regimes since the Last Glacial Maximum: an assessment based on a global synthesis and analysis of charcoal data. *Climate Dyn.* **30**, 887–907 (2008).
- Guenther, A. *et al.* A global model of natural volatile organic compound emissions. *J. Geophys. Res.* **100**, 8873–8892 (1995).
- Prather, M., Holmes, C. & Hsu, J. Reactive greenhouse gas scenarios: systematic exploration of uncertainties and the role of atmospheric chemistry. *Geophys. Res. Lett.* **39**, L09803 (2012).
- Naik, V. *et al.* Preindustrial to present-day changes in tropospheric hydroxyl radical and methane lifetime from the atmospheric chemistry and climate model intercomparison project (ACCMIP). *Atmos. Chem. Phys.* **13**, 5277–5298 (2013).
- Prather, M. *et al.* in *Climate Change 2001: The Physical Science Basis: Contribution of Working Group I to the Third Assessment Report of the Intergovernmental Panel on Climate Change* (eds Houghton, J. *et al.*) (Cambridge University Press, Cambridge, UK & NY, USA, 2001).
- Law, K. & Pyle, J. Modeling trace gas budgets in the troposphere 1. Ozone and odd nitrogen. *J. Geophys. Res.* **98**, 18377–18400 (1993).
- Wild, O., Zhu, X. & Prather, M. Fast-J: accurate simulation of in- and below-cloud photolysis in tropospheric chemical models. *Atmos. Chem. Phys.* **37**, 245–282 (2000).
- Sapart, C. *et al.* Natural and anthropogenic variations in methane sources during the past two millennia. *Nature* **490**, 1723–1726 (2012).
- Kohfeld, K. & Ridgwell, A. in *Surface Ocean-Lower Atmosphere Processes* (eds LeQuere, C. & Saltzman, E.) Geophysical Research Series 187, 251–286 (AGU, 2009).
- Möller, L. *et al.* Independent variations of CH_4 emissions and isotopic composition over the past 160,000 years. *Nat. Geosci.* **6**, 885–891 (2013).
- Tan, Z., Zhuang, Q. & Water-Anthony, K. Modeling methane emissions from arctic lakes: model development and site-level study. *J. Adv. Model Earth Syst.* **7** (2015).
- Ringeval, B. *et al.* Methane emissions from floodplains in the Amazon Basin: challenges in developing a process-based model for global applications. *Biogeosciences* **11**, 1519–1558 (2014).
- Fan, Y. & Miguez-Macho, G. A simple hydrologic framework for simulating wetlands in climate and earth system models. *Clim. Dyn.* **37**, 253–278 (2011).
- Boardman, C., Gauci, V., Fox, A., Blake, S. & Beerling, D. Reduction of the temperature sensitivity of minerotrophic fen methane emissions by simulated glacial atmospheric carbon dioxide starvation. *J. Geophys. Res.* **118**, 1–9 (2013).
- Boardman, C., Gauci, V., Watson, J., Blake, S. & Beerling, D. Contrasting wetland CH_4 emission responses to simulated glacial atmospheric CO_2 in temperate bogs and fens. *New Phytol.* **191**, 898–911 (2012).
- Williams, E. Lightning and climate: A review. *Atmos. Res.* **76**, 272–287 (2005).
- Hopcroft, P., Valdes, P., Woodward, S. & Joshi, M. Last glacial maximum radiative forcing from mineral dust aerosols in an Earth System model. *J. Geophys. Res.* **120**, 8186–8205 (2015).

47. Kohfeld, K. & Harrison, S. DIRTMAP: the geological record of dust. *Earth Sci. Rev.* **54**, 114 (2001).
48. Yuan, T., Remer, L., Pickering, K. & Yu, H. Observational evidence for aerosol enhancement of lightning activity and convective invigoration. *Geophys. Res. Lett.* **38**, L04701 (2011).
49. Monson, R., Grote, R., Niinemets, U. & Schnitzler, J.-P. Modeling the isoprene emission rate from leaves. *New Phytol.* **195**, 541–559 (2012).
50. Morfopoulos, C. *et al.* A unifying conceptual model for the environmental responses of isoprene emissions from plants. *Ann. Bot. (Lond.)* **112**, 1223–1238 (2013).
51. Hopcroft, P. & Valdes, P. Last Glacial Maximum constraints on the Earth System model HadGEM2-ES. *Clim. Dyn.* **45**, 1657–1672 (2015).
52. Walters, D. *et al.* The Met Office Unified Model Global Atmosphere 4.0 and JULES Global Land 4.0 configurations. *Geosci. Model Dev.* **7**, 361–386 (2014).
53. Nowack, P. *et al.* A large ozone-circulation feedback and its implications for global warming assessments. *Nat. Clim. Change* **5**, 41–45 (2015).
54. Marsh, D., Lamarque, J.-F., Conley, A. & Polvani, L. Stratospheric ozone chemistry feedbacks are not critical for the determination of climate sensitivity in CESM1(WACCM). *Geophys. Res. Lett.* **43**, 3928–3934 (2016).
55. Baumgartner, M. *et al.* NGRIP CH₄ concentration from 120 to 10 kyr before present and its relation to a 15N temperature reconstruction from the same ice core. *Clim. Past* **10**, 903–920 (2014).
56. Carslaw, K. *et al.* Large contribution of natural aerosols to uncertainty in indirect forcing. *Nature* **503**, 67–71 (2013).
57. Cox, P. Description of the TRIFFID dynamic global vegetation model. *Hadley Centre, Met Office, Technical Note* **24** (2001).
58. Jones, C. *et al.* The HadGEM2-ES implementation of CMIP5 centennial simulations. *Geosci. Model Dev.* **4**, 543–570 (2011).
59. Pöschl, U. *et al.* Development and intercomparison of condensed isoprene oxidation mechanisms for global atmospheric modeling. *J. Atmos. Chem.* **37**, 29–52 (2000).
60. Price, C. & Rind, D. A simple lightning parameterization for calculating global lightning distributions. *J. Geophys. Res.* **92**, 9919–9933 (1992).
61. Price, C. & Rind, D. Modeling global lightning distributions in a general circulation model. *Mon. Weather Rev.* **122**, 1930–1939 (1994).
62. Hopcroft, P. & Valdes, P. How well do simulated last glacial maximum tropical temperatures constrain equilibrium climate sensitivity? *Geophys. Res. Lett.* **42**, 5533–5539 (2015).
63. Gedney, N. & Cox, P. The sensitivity of global climate model simulations to the representation of soil moisture heterogeneity. *J. Hydrometeorol.* **4**, 1265–1275 (2003).
64. Best, M. *et al.* The joint UK land environment simulator (JULES), model description—part 1: energy and water fluxes. *Geosci. Model Dev.* **4**, 677–699 (2011).
65. Matthews, T., Dadson, S., Lehner, B., Abele, S. & Gedney, N. A high-resolution global dataset of topographic index values for use in large-scale hydrological modelling. *Hydrol. Earth Syst. Sci.* **19**, 91–104 (2015).
66. Armante, C. & Eakins, B. ETOPO1 1 arc-minute global relief model; procedures, data sources and analysis. Tech. Rep., NOAA Technical Memorandum, NESDIS NGDC-24 (2009).
67. Lamarque, J.-F. *et al.* Historical (1850–2000) gridded anthropogenic and biomass burning emissions of reactive gases and aerosols: methodology and application. *Atmos. Chem. Phys.* **10**, 7017–7039 (2010).
68. Kaplan, J., Pfeiffer, M., Kolen, J. & Davis, B. Large scale anthropogenic reduction of forest cover in last glacial maximum Europe. *PLoS ONE* **11**, e0166726 (2016).
69. Sanderson, M. Biomass of termites and their emission of methane and carbon dioxide: A global database. *Glob. Biogeochem. Cyc.* **10**, 543–557 (1996).
70. Curry, C. Modeling the soil consumption of atmospheric methane at the global scale. *Glob. Biogeochem. Cyc.* **21**, GB4012 (2007).
71. Wieder, W., Boehnert, J., Bonan, G. & Langseth, M. RegridDED Harmonized World Soil Database v1.2. Data set (2014).
72. Pacifico, F. *et al.* Evaluation of a photosynthesis-based biogenic isoprene emission scheme in JULES and simulation of isoprene emissions under present-day climate conditions. *Atmos. Chem. Phys.* **11**, 4371–4389 (2011).
73. Pacifico, F., Folberth, G., Jones, C., Harrison, S. & Collins, W. Sensitivity of biogenic isoprene emissions to past, present, and future environmental conditions and implications for atmospheric chemistry. *J. Geophys. Res.* **117**, D22302 (2012).
74. Yienger, J. & Levy, II H. Empirical model of global soil-biogenic NO_x emissions. *J. Geophys. Res.* **100**, 11447–11464 (1995).
75. Kageyama, M. *et al.* Mid-Holocene and last glacial maximum climate simulations with the IPSL model: Part I: comparing IPSL-CM5A to IPSL-CM4. *Clim. Dyn.* **40**, 2447–2468 (2013).
76. Kleinen, T., Brovkin, V. & Schuldt, R. A dynamic model of wetland extent and peat accumulation: results for the Holocene. *Biogeosciences* **9**, 235–248 (2012).
77. Spahni, R. *et al.* Constraining global methane emissions and uptake by ecosystems. *Biogeosciences* **8**, 1643–1665 (2011).
78. Voulgarakis, A. *et al.* Analysis of present day and future OH and methane lifetime in the ACCMIP simulations. *Atmos. Chem. Phys.* **13**, 2563–2587 (2013).

Acknowledgements

HadGEM2-ES simulations were performed on the UK national computing facility ARCHER (<http://www.archer.ac.uk>). We thank Garry Hayman for configuring the Fast-J photolysis scheme with the TropIsop chemistry scheme. We thank Lee Murray for providing modelled O₃ fields and the IPSL group for making their PMIP3 model output available via CMIP5 and ESGF. P.O.H. was funded by the NERC project NE/I010912/1. The development of HadGEM2-ES and F.M.O'C. were supported by the Joint UK BEIS/Defra Met Office Hadley Centre Climate Programme (GA01101). J.O.K. was supported by the European Research Council (313797, COEVOLVE).

Author contributions

P.O.H. and P.J.V. designed the experiments with input from F.M.O'C. J.O.K. designed and carried out fire simulations, all other model simulations performed by P.O.H. P.O.H. wrote the paper with input from P.J.V., F.M.O'C., J.O.K. and D.J.B.

Additional information

Supplementary Information accompanies this paper at <http://www.nature.com/naturecommunications>

Competing financial interests: The authors declare no competing financial interests.

Reprints and permission information is available online at <http://npg.nature.com/reprintsandpermissions/>

How to cite this article: Hopcroft, P. O. *et al.* Understanding the glacial methane cycle. *Nat. Commun.* **8**, 14383 doi: 10.1038/ncomms14383 (2017).

Publisher's note: Springer Nature remains neutral with regard to jurisdictional claims in published maps and institutional affiliations.



This work is licensed under a Creative Commons Attribution 4.0 International License. The images or other third party material in this article are included in the article's Creative Commons license, unless indicated otherwise in the credit line; if the material is not included under the Creative Commons license, users will need to obtain permission from the license holder to reproduce the material. To view a copy of this license, visit <http://creativecommons.org/licenses/by/4.0/>

© The Author(s) 2017